Erniggliite (Tl2SnAs2S6), a new mineral from Lengenbach, Binntal (Switzerland) : description and crystal structure determination based on data from synchrotron radiation

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Erniggliite ($\text{TI}_2\text{SnAs}_2\text{S}_6$), a new mineral from Lengenbach, Binntal (Switzerland): description and ciystal structure determination based on data from synchrotron radiation

Dedicated to Prof. Ernst Niggli on the occasion of his 75th anniversary

by Stefan Graeser^{1,2}, Hans Schwander², Reinhard Wulf³ and Andreas Edenharter³

Abstract

Erniggliite, $Tl_2SnAs_2S_6$, is a new mineral species from the famous sulfosalt locality Lengenbach, Binntal (Switzerland). It occurs in small cavities in a dolomitic rock of Triassic age. Associated are a large number of various, mainly Tl-bearing As-sulfosalts such as hutchinsonite, hatchite, wallisite, lorandite, bernardite, abundant realgar and orpiment and two additional new minerals, edenharterite and stalderite. The mineral is characterized by its conspicuous hexagonal habit and an uncommon micaceous cleavage parallel to (0001). Individual crystals are very small and do not exceed 0.5 mm. Chemical composition (electron microprobe, mean of 3 analyses) is: T1 47.91, Sn 13.55, As 17.01, S 22.20, total 100.67 wt% leading to the mineral formula $Tl_2SnAs_2S_6$. The mineral has a steel-grey colour, in fresh cleavage fragments it occurs shiny black, the streak is reddish-black. Microhardness is 47.7 kg/mm² for 10 g load (range 46–49), corresponding to a Mohs hardness of about 2–3, $D_{calc} = 5.24$ g/cm³. Single crystal studies yielded a trigonal cell with a = 6.680 (3), c = 7.164 (9) Å, V = 276.9 (3) Å³, space group is P3 and Z = 1. The strongest lines in the X-ray powder pattern are (d_{obs}, I_{obs}, hkl) : 3.343 (100)(110), 3.029 (63)(111), 3.060 (50)(102), 2.679 (48)(201), 4.510 (40)(011), 1.866 (38)(212), 2.25 (37)(202), 1.930 (36)(030).

Single crystal diffraction measurements were carried out with synchrotron radiation ($\lambda = 1.1573$ Å) on the 5circle diffractometer at HASYLAB, Hamburg. The structure was solved by an Automated Patterson Interpretation (SHELXS 92) using 448 unique reflections. The refinement of positional parameters and anisotropic temperature factors under consideration of anomalous dispersion correction terms and isotropic extinction lead to ^a final weighted R value of 0.0253 for all data. In this structure, As builds ^a trigonal pyramid with three sulfur atoms at the base and As at the top of the polyhedron which is common in As-bearing sulfosalts. However, T1 and Sn form two new coordination polyhedra with sulfur, not known so far from sulfosalts and Tl-Sn-sulfides. The Tl-polyhedron can be described as an antiprism formed by ^a slightly deformed hexagonal and trigonal base with the central atom T1 shifted towards the hexagonal base. Moreover, ^a remarkably short Tl-As distance of 3.259 Â is observed which is slightly shorter than the shortest Tl-S bonding of 3.268 A in this structure. The Sn-polyhedron is a nearly perfect trigonal antiprism with six equal Sn-S distances of 2.5614 A. The structure is based on layers // (0001) which are formed by ^a three dimensional network of Tl-polyhedra and As-pyramids. These layers are only connected by the Sn-S antiprisms, explaining the excellent cleavage of the crystals.

The name is for Ernst Niggli, Professor emeritus in mineralogy and petrography at the University of Berne, and president for many years of the "Arbeitsgemeinschaft Lengenbach".

Keywords: Erniggliite, new mineral, Tl-Sn-As-sulfosalt, microprobe analyses, synchrotron radiation, Lengenbach, Binntal, Switzerland.

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1. Introduction

The locality Lengenbach, Binntal, situated in an outcrop of Triassic dolomite of the Penninic Mte. Leone nappe, is known for its special sulfide and sulfosalt minerals since more than 200 years. Af^a break of almost ⁵⁰ years, from the beginning of World War I until 1958 during which the traction of the special minerals had come to a complete standstill, ^a syndicate of Swiss museums and university institutions ("Arbeitsgemeinschaft Lengenbach", AGL) resuscitated the exploitation works. For more than 20 years this activity had been relatively successful, but then, in the early eighties, the extraction of interesting minerals came more and more problematic. This was partly due to technical reasons, but even more to mineralogical considerations, which made further work in the classical quarry very uncertain. On the one hand, the excavation of minerals at one small site over almost 200 years had led to ^a deep hole with perpendicular walls up to ²⁰ meters high that endangered any activity at its bottom. On the other hand, the number of interesting erals gradually decreased with increasing depth. Finally, two alternatives arose for the syndicate – the question of either giving up any exploitation or starting at ^a new site. Exactly at this critical moment, in 1985, ^a huge loose block fell into the old quarry which, when split up, turned out to be full of highly interesting minerals. Within less than two years, the study of this material duced three new mineral species, all of them representing new thallium-arsenic-sulfosalts: edenharterite, $TIPbAs₃S₆$ (GRAESER and SCHWANDER, 1992), stalderite, $TICu(Zn,Fe,Hg)$ ₂As₂S₆ (GRAE-

Tab. ¹ Physical and optical properties of erniggliite from Lengenbach.

 $^{2})$ Edenharterite, TlPb $\mathrm{As}_{3}\mathrm{S}_{6}$, Lengenbach (Gr $\mathrm{As}_{5}\mathrm{S}_{8}$ and SCHWANDER, 1992)

ser et al., in prep.) and erniggliite, $Tl_2SnAs_2S_6$ (this paper). The discovery of the three new minerals clearly favoured the syndicate's decision to tinue work at a higher level in an easterly direction instead of giving up any mineral extraction at

Fig. ¹ Erniggliite, Lengenbach. On the SEM picture the (pseudo-) hexagonal habit of the mineral is easily recognizable. Forms are: pinacoid (0001), hexagonal prism and rhombohedral faces. The surface is rough and uneven (picture: SEM laboratory, University of Basel).

Fig. 2 Erniggliite, Sem picture under stronger magnification documenting the extremely good cleavage parallel to $\{0001\}$ (SEM laboratory, University of Basel).

	A	B	C	D	Е		
Tl	47.78	47.50	48.44	47.91	47.00 wt %		
Sn	13.52	13.56	13.57	13.55	13.65		
As	16.97	17.05	17.02	17.01	17.23		
S	22.16	22.36	22.07	22.20	22.12		
	Total 100.43	100.47	101.10	100.67	100.00 wt\%		
Structural formula (based on 11 atoms):							
$T1^{1+}$	2.03	2.01	2.06	2.03	2.0		
Sn^{4+}	0.99	0.99	0.99	0.99	1.0		
As^{3+}	1.97	1.97	1.97	1.97	2.0		
S	6.01	6.03	5.98	6.01	6.0		
A, B, C analyses of individual grains mean of 3 analyses D theoretical composition for $Tl_2SnAs_2S_6$ Е							

Tab. 2 Electron microprobe analyses of erniggliite from Lengenbach, Binntal.

Tab. 3 X-ray diffraction data for erniggliite from Lengenbach.

Erniggliite was detected when looking through the share of the Natural History Museum Basel from the exploitation period 1985 at Lengenbach for additional material of the previously mined new mineral stalderite. No more material of this mineral could be found yet, but another

0001 0111 $10₁₁$ $10₁₀$

Fig. 3 Erniggliite, idealized computer drawing (SHAPE), showing the common habit of the mineral.

Debye-Scherrer camera, 90 mm diameter, FeKaradiation, intensities determined densitometrically.

Cell parameters refined from the powder data.

unknown mineral was found in sample number L 18,393 that turned out to be the new mineral erniggliite. It is ^a remarkable fact that another new mineral, edenharterite, was also found in this same sample (L 18,393).

Name and properties of the new mineral were submitted to the "Commission on New Minerals and Mineral Names, I.M.A." for consideration. Under the number 87-025 the proposal for the new mineral was accepted unanimously by the commission.

2. Description

2.1. PHYSICAL AND OPTICAL PROPERTIES

When looking through the samples of the 1985 period, ^a mineral with ^a conspicuous hexagonal morphology attracted our attention, because, up to that moment, no sulfosalt mineral with such a morphology was known from Lengenbach. liminary Gandolfi-studies confirmed this assump-

Tab. 4 Summary of synchrotron data-collection parameters, structure solution and refinement of erniggliite, $Tl_2SnAs_2S_6.$

tion, as the Gandolfi diagram was not identifiable with that of any known mineral species.

Erniggliite occurs in druses (solution cavities) and fissures in the dolomite. The mineral forms euhedral crystals of short prismatic, clearly agonal habit less than 1 mm in length. The surface of the crystals is rough, especially when observed under higher magnification. It reveals a conspicuously good, micaceous cleavage perpendicular to the c-axis, very uncommon to other Lengenbach sulfosalts. The colour is grey to black on its surface but shiny black on fresh cleavage planes. The mineral is opaque, even in thin cleavage lamellae; these lamellae are flexible, comparable with thin

crystals of the mineral lengenbachite. The streak colour is reddish-black, very similar to that of stalderite, but clearly distinctive from edenharterite. The hardness, measured as Vicker's hardness numbers in the polished section gave values between 46—49 kg/mm2, according to a Mohs hardness of 2-3.

Due to the micaceous structure of the mineral, optical properties could be studied only in tions parallel to $\{0001\}$. Under the ore microscope, therefore, the mineral appeared completely isotropic. The colour of the mineral in reflected light is (greyish-) white, without bireflectance or pleochroism phenomena; internal reflections

	$\mathbf x$	v	z		$U_{\text{equ}}[\text{\AA}^2]$		
Sn	0.0000	0.0000	0.0000		0.0103(3)		
T1 As	0.3333 0.3333	0.6667 0.6667	0.3823(1) 0.9266(1)		0.0377(3)		
S	0.2051(2)	0.3346(2)	0.7684(2)	0.0161(4) 0.0196(7)			
	U_{11}	U_{22}	U_{33}	$\rm{U_{23}}$	\mathbf{U}_{13}	U_{12}	
Sn	0.0077(3)	0.0077(3)	0.0156(3)	0.000	0.000	0.0039(2)	
Tl	0.0461(3)	0.0461(3)	0.0208(3)	0.000	0.000	0.0231(2)	
As	0.0163(4)	0.0163(4)	0.0157(4)	0.000	0.000	0.0081(2)	
S	0.0269(7)	0.0102(6)	0.0180(5)	0.0015(4)	0.0056(5)	0.0065(5)	

Tab. 5 Final fractional atomic coordinates and anisotropic thermal parameters U_{ii} of erniggliite (Tl₂SnAs_{2S6}) from synchroton data ($\lambda = 1.1573 \text{ Å}$). T_{hkl} = exp[-2 π^2 ·(U₁₁h²a*²+U₂₂k²b*²+U₃₃l²c*2+2U₂₃klb*c*+2U₁₃hla*c*+2U₁₂hka*b*)]

could not be observed. All the properties given in table ¹ were obtained from such sections as no sections parallel to the c-axis were preparable.

With one of the tiny crystals (about 0.5 mm long) we tried to identify the indices of the face forms by measurement with an optical goniome-Unfortunately the rough quality of the faces prevented the detection of clear reflexion signals. Nevertheless, an indistinct light concentration dicated a value of about 54° for the angle ρ of the rhombohedrons, and a similar value was obtained by measurement of the angle between the pinacoid {0001} and the rhombohedral faces der the microscope. The calculated angle ρ for the form {1011} of erniggliite is 51.1°, so, with some probability, the rhombohedral faces have the indices $\{10\overline{1}1\}$ and $\{01\overline{1}1\}$, respectively (see Fig. 3).

2.2. CHEMICAL COMPOSITION

Preliminary investigations have been carried out by SEM techniques (Cambridge Stereoscan Mark 2A, equipped with EDS system). The SEM tures display the hexagonal habit of the crystals and their rough surface and, moreover, the spicuous micaceous cleavage of erniggliite is clearly documented by the pictures under stronger magnification (see Figs 1, 2). The EDS diagram unequivocally indicated the composition of the major elements Tl, Sn, As, S. Among the Tlsulfosalts known so far, no compound between Tl, Sn, As could be found; even from the chemical composition it became obvious that the mineral must represent ^a new mineral species. So the next step was to obtain a quantitative chemical deter-

mination. Three chemical analyses were carried out on individual grains by means of an electron microprobe (ARL instrument, SEMQ with EDSsystem TN 2000, ZAF program). The following materials were used as standards: $TIAsS₂$ (Tl, As, S), galena (Pb), and pure elements (Sn, Pb). The results of the analyses are shown in table 2. From the analytical data a mineral formula with the expression $Tl_2SnAs_2S_6$ could be calculated. Erniggliite represents the first Lengenbach sulfosalt with Sn as ^a major element and, moreover, it is the first and only Tl-Sn-sulfosalt known so far.

When it became obvious that the mineral under investigation represented a new mineral, we informed Prof. G.H. Moh (Heidelberg), a keen specialist in Tl-sulfosalts, about the preliminary qualitative composition (Tl-Sn-As-S) of the known mineral. One of his coworkers, N. Wang, started experimental studies with this information. He succeeded in synthesizing a powdery sulfosalt phase from these elements that, by an X -ray powder diffraction study, turned out to be lutely identical with our new mineral. The results of this experimental study will be published where (WANG and GRAESER, in prep.).

2.3. X-RAY DIFFRACTION STUDY

First diffraction experiments were carried out by Gandolfi techniques. They yielded a simple and characteristic pattern which was not identifiable with that of any known mineral species. In order to eliminate any (non-) orientation effects by the Gandolfi techniques, a small fragment of the cryswas sacrificed for ^a normal Debye-Scherrer powder pattern (Tab. 3) which turned out to be

distances (\AA)				angles (°)		
3.2594(4) $T1 -$ As Tl 3.2681(6) $S_{4,5,6}$ 3.3695(5)	3x $3\times$	As	T1 $\overline{}$ T1 Tl	$S_{4,5,6}$ $S_{1,2,3}$ $S'_{4,5,6}$	70.767(5) 144.96(1) 73.305(4)	3x 3x $3\times$
$S_{1,2,3}$ $S'_{4,5,6}$ 3.7473(7) mean Tl-S 3.4616	$3\times$	S_4 S_5 $S_{4,5,6}$ $S_{4,5,6}$ $S_{4,5,6}$ $S_{4,5,6}$ $S_{4,5,6}$ $S_{4,5,6}$ S_1 S ₂ $S_{1,2,3}$ $S_{1,2,3}$ $S_{1,2,3}$ S_4^- S_5	Tl Tl Tl TI Tl ÷, Tl $\overline{}$ Tl $\overline{}$ Tl - Tl ÷ Tl - Tl \rightarrow Tl $\mathbf T\mathbf l$ Tl $TI -$	$S_{5,6}$ S_6 $S_{1,2,3}$ $S_{2,3,1}$ $S_{3,2,1}$ $S'_{4,5,6}$ $S'_{5,4,6}$ $S'_{6,5,4}$ $S_{2,3}$ S_3 $S'_{4,5,6}$ $S_{5,6,4}$ $S_{6,4,5}^{\dagger}$ $S'_{5,6}$ S_6^r	109.709(6) 109.709(6) 101.395(3) 80.825(3) 140.45(2) 144.07 (2) 56.56(1) 57.14(1) 59.626(8) 59.626(8) 107.786(5) 74.204(4) 132.25(1) 112.098(5) 112.098(5)	$2\times$ $3\times$ $3\times$ 3x 3x $3\times$ $3\times$ $2\times$ $3\times$ $3\times$ 3x $2\times$
$Sn -$ 2.5614(6) $S_{1,2,3}$ 2.5614(6) $S_{4,5,6}$	$.3\times$ $3\times$	S_1 S_2 S_4 S_5	$Sn -$ $Sn -$ $Sn -$ $Sn -$	$S_{2,3}$ S_3 $S_{5,6}$ S_6	82.607(3) 82.607(3) 97.393(3) 97.393(3)	$2\times$ $2 \times$
2.2421(8) As $ S_{1,2,3}$ $\chi \propto e^{-\lambda}$	$3\times$	\mathbf{S}_1 S ₂	$As -$ $As -$	$S_{2,3}$ S_3	96.691(3) 96.691(3)	$2\times$
\mathcal{C} \mathcal{P}^{\bullet} : Symmetry code	e Roma					
$\lambda \neq$ $\ddot{1}$: z X, y, 4: $-X,$ $-y,$ $-z$	2: 5:	$-y,$ x-y, у, y-x,	$\mathbf{Z}% _{M_{1},M_{2}}^{\alpha,\beta}(\mathbf{Z}_{M_{1},M_{2}}^{\alpha,\beta}(\mathbf{Z}% _{M_{1},M_{2}}^{\beta,\beta}))$ $-Z$		3: y-x, 6: x-y,	\mathbf{Z} -x, $-\mathbf{Z}$ x,

Tab. 6 Interatomic distances and angles in erniggliite (standard deviations in brackets).

practically identical with the Gandolfi pattern, with minor shifts of the intensities. Consequently, the mineral was studied by a single crystal investigation with Weissenberg and Precession techniques. Due to the micaceous cleavage and the flexibility of the mineral, it was quite difficult to separate suitable single crystal fragments. From the single crystal study a cell with hexagonal metry was derived, primarily, yet the subsequent structure determination clearly proved the trigonal symmetry of the mineral.

2.4. OCCURRENCE, GEOCHEMICAL CONSIDERATIONS

Erniggliite is not only the first Sn-bearing sulfosalt mineral of the Lengenbach occurrence, it seems to represent the first Tl-Sn-As-sulfosalt scribed to date. In the long history of more than 250 years of mineralogical research in Lengenbach, no mineral with Sn as a major element has been detected among the wide spectrum of its minerals. As a minor element, however, Sn could be determined in the course of ^a trace element study of some sulfide minerals from the dolomitic rocks in Binntal; galena and sphalerite as well may contain the element Sn in concentrations up to 400 ppm (Graeser, 1969). In various studies on the formation of the remarkable As-sulfosalt eralization at Lengenbach, we postulated the hydrothermal transport of several elements like As, Cu, Tl, etc. from outside (by remobilization of Cu-As-ore concentrations in the gneisses in the south) into the dolomitic rocks where they formed the sulfosalt minerals by reaction with the ERNIGGLIITE, A NEW MINERAL FROM LENGENBACH 299

Fig. 4 Projection of the structure of erniggliite $//$ c-axis (the unit cell is marked with broken lines).

minerals preexistent in the dolomite, like galena, pyrite, sphalerite, etc. (Graeser 1965, 1975). On their way through the gneisses towards north, these hydrothermal solutions gradually cooled, became oversaturated, and were partly precipitated, leading to a very uncommon mineralization within the gneisses. One of these special minerals in the gneisses, asbecasite, a Be-Ca-As-oxide scribed as a new mineral species by GRAESER, 1966), showed an unexpected content of 2.1 wt% $SnO₂$. Obviously, the element Sn too, was transported into the dolomites together with As, Cu, Tl, etc. and, under extremely special conditions, even led to the formation of sulfosalts.with Sn as a major element in the dolomite.

Erniggliite, in its characteristic habit, represents the only Lengenbach sulfosalt mineral with unequivocally (pseudo-) hexagonal symmetry and together with its conspicuous excellent cleavage along (0001), may not be mistaken for any other Lengenbach mineral. Therefore, it is highly probable that erniggliite occurred in the old clas-Lengenbach quarry - such a characteristical mineral would not have escaped the keen attention of the old mineralogists! The mineral was found together with the two other new minerals edenharterite, TlPbAs₃S₆ (GRAESER and SCHWAN-DER, 1992) and stalderite, $TICu(Zn,Fe,Hg)$ ₂As₂S₆ (Graeser et al., in prep.) in ^a loose block from the upmost border of the classical Lengenbach quarry. So it presumably originates from a considerably higher level than had ever been worked for the extraction of minerals. It is ^a strange fact that in the very restricted area of the mineralized dolomitic rock in Lengenbach, geochemical factors change rapidly within a few meters. The special zone of the dolomite containing the Tl-rich minerals obviously could not be found in the old quarry

Fig. 5 Projection of the structure of erniggliite $// a_2^*$ -axis.

because it was eroded by the Lengenbach stream 3. Experimental long before. For details of the differences tween the old and the new working sites in Len- 3.1. USE OF SYNCHROTRON RADIATION genbach, see also the description of edenharterite

minerals. Edenharterite has been found in more than 100 clearly identified samples since its disall. The Tl-Sn-As-sulfosalt erniggliite seems to

(GRAESER and SCHWANDER, 1992). Synchrotron radiation results from the accelera-
Meanwhile, 5–6 years have passed since the tion of relativistic electrons in storage rings, which tion of relativistic electrons in storage rings, which
were primarily used for high energy collision exfirst indications of the three new Lengenbach were primarily used for high energy collision ex-
minerals. Edenharterite has been found in more periments, and was regarded for a long time only as an energy consuming undesired byproduct.
Meanwhile, due to the excellent radiation propercovery. In the case of stalderite, about 10 samples Meanwhile, due to the excellent radiation proper-
have been collected, but for erniggliite only two ties, e.g. the high intensity at a wide spectral range have been collected, but for erniggliite only two ties, e.g. the high intensity at a wide spectral range samples exist, with about 6–8 minute crystals, in up to the hard X-ray range, these storage rings samples exist, with about 6-8 minute crystals, in up to the hard X-ray range, these storage rings all. The Tl-Sn-As-sulfosalt erniggliite seems to have been improved and optimized especially for represent by far the rarest member of the three the generation of synchrotron radiation. Hereby new Tl-minerals from Lengenbach. the high intensity wavelength range depends in

Fig. 6 Me-coordination polyhedra. a) TlS₉-polyhedron viewed along the c-axis; b) TlS₉-polyhedron rotated 50° around the a_2 -axis; c) SnS_6 trigonal antiprism; d) AsS_3 trigonal pyramid.

each case on the characteristics of the storage rings. Consequently, worldwide several laboratohave been built up with specific experimental stations. For crystallographers and mineralogists the development of high precision single crystal diffractometry and powder diffraction are of special interest. The main advantages of using synchrotron radiation in single crystal work are:

1) Variabel wavelength \rightarrow the wavelength can be adapted to the problem studied (e.g. use of anomalous dispersion effects).

2) High intensity \rightarrow smaller crystals can be used to reduce problems arising from faulty extinction

and absorption correction. Additionally, even very weak reflections are measured with ^a high significance level.

3) High degree of polarization \rightarrow reflections with a high 2Θ -angle can be measured significantly.

To take advantage of these features we decided to use synchrotron radiation in case of the new mineral erniggliite. Our aim was, to obtain an excellent data set for a precise structure determination with low standard deviation of the atomic parameters in order to get exact structural data for the systematics of Tl-Sn sulfosalts. Since

 $\hat{\boldsymbol{\theta}}$

 \mathbf{R}

 $Tl_2SnAs_2S_6$ shows an excellent cleavage parallel (0001), it was necessary, to use ^a crystal as small as possible, to avoid micro-twinning and splitting of the reflection profiles.

Our single crystal X -ray diffraction measurement was carried out at the five circle diffractometer at HASYLAB (Hamburger Synchrotron bor) at DESY (Deutsches Elektronen-Synchrotron). The main components of this experimental station for single crystal investigations, which was designed and mounted by the synchrotron workgroup of Prof. Dr. Kupäk from the Institute of mineralogy and crystallography, (KUPČIK et al., 1983, 1986) are:

a) commercial four circle diffractometer (Stoe Stadi4) which is modified and adapted for synchrotron radiation use and mounted on a fifth circle for special applications,

b) fixed exit monochromator with two independent germanium [111] single crystals,

c) polarization/intensity monitor for data duction and beam stability control,

d) beam position monitor.

The main specifications of crystallographic interest are:

- a) angular resolution: 0.001°
- b) wavelength range: 0.3-2.2 A
- c) wavelength resolution at ¹ Â:
- $\Delta\lambda/\lambda \approx 3 \times 10^{-4}$
- d) degree of polarization $\approx 85-95\%$

For the actual diffraction measurements a tiny crystal of good quality with dimensions of 0.173 \times 0.154×0.116 mm³ was used.

From 23 reflections in the 2©-range of 21.8- 81.8° the lattice constants were determined by least-squares refinement to be $a = 6.678(2)$, c = 7.160(2), $\alpha = 90^{\circ}$, $\beta = 90^{\circ}$ and $\gamma = 120^{\circ}$. With $Z = 1$ a calculated density of $\rho = 5.233$ g/cm³ results. These values are in good agreement with those evaluated from the powder measurements tioned above.

The data collection was carried out at 1.1573 Â using a modified measuring procedure (WENDschuh-JOSTIES and Wulf, 1989) which takes into account the special properties of synchrotron diation.

In the range of $(\sin\Theta)/\lambda \leq 0.66$ Å⁻¹ 1589 reflections were measured in the ω step scan mode with ^a step width of 0.01°. The intensity data were scaled by monitor and standard intensities and corrected for Lorentz and polarization factors ing the synchrotron data reduction program Synred (WENDSCHUH-JOSTIES and WULF, 1984). Additionally, an empirical absorption correction based on ψ -scans was carried out (XEMP, SHELDRICK, 1976). The linear absorption coefficient at $\lambda =$

1.1573 Å is calculated to μ = 394.61 cm⁻¹ resulting to a mean μ R-value of 2.91. Table 4 summarizes the parameters and results of the data collection and the subsequent structure solution and refinement. Analysing the intensity data for systematic absences and internal consistency of equivalent reflections the spacegroups P3 and P3 remained possible. Finally, from structure refinement the spacegroup P3 turned out to be the correct one.

3.2. STRUCTURE SOLUTION AND REFINEMENT

Whereas direct methods failed, the structure solution succeeded by use of an automated supersharp Patterson superposition method, implemented in the new structure solution package SHELXS-92 (SHELDRICK, 1992) which was available at that time as a pre-version. For the structure refinement we used the programm SHELXL-92 $(SHELDRICK, 1992)$, where the least-squares refinement is based on $F²$. Taking into account anomalous dispersion correction terms, the refinement of atomic parameters, anisotropic temperature factors and isotropic extinction led to an unweigthed $R(F)$ -value of 0.0242 for observed (445) and 0.0253 for all reflections (448). It is important to notice, that as a result of the high intensity and collimation of synchrotron radiation from ^a total of 448 unique reflections ⁴⁴⁵ were measured with $F^2 > 2 \sigma(F^2)$. Additionally the standard deviations of atomic and thermal parameters are remarkable small. A refinement of the occupancy factors confirm the ideal formula of $Tl_2SnAs_2S_6$.

According to the classification of sulfosalts as proposed by Nowacki (1969), erniggliite with $S:As = \varphi = 3$ belongs to the group II α_1 with isolated BS_3 -pyramids, which is a common group in As-sulfosalts.

The final atomic coordinates and thermal rameters including e.s.d.'s are summarized in table 5, while the interatomic distances and angles are listed in table 6. For completeness an abbreviated F_o/F_c list is given in table 7.

3.3. DESCRIPTION OF THE STRUCTURE

The arrangements of the atoms projected //c- and // a_2^* -axis are shown in figures 4 and 5 respectively. In figures 6 a-d the Me-polyhedra are presented.

Analysing the structure of erniggliite we found two new (Me-S)-polyhedra, a TlS₉- and a $SnS₆$ polyhedron not known so far from sulfosalts and (Tl-Sn)-sulfides. The coordination polyhedron of T1 may be described in two different ways. If we

assume, that Tl-S distances up to 3.75 Å are regarded as bonds (mean value 3.4616 Å), the resulting polyhedron is an antiprism with different bases. However, if we only consider the six short-Tl-S distances as bonds (mean value 3.3188 Â) a distorted trigonal antiprism with additionally three longer distances is formed. Since these longbonds are within the plane of one trigonal base, we prefer the first description and propose for this new kind of environment the name "hex-tri" tiprism. In this surrounding the Tl atom is coordiby nine sulfur atoms with distances in the range of 3.2681 to 3.7473 Â (Tab. 6) in form of ^a slightly distorted antiprism with ^a hexagonal and a trigonal base (Figs $6a, b$). The Tl atom is shifted from the center of the polyhedron along the 3-fold rotation axis towards the hexagonal base. This base is built by those three sulfur atoms with the shortest $(3 \times 3.2681$ Å) and those three sulfur atoms with the longest Tl-S bonds $(3 \times 3.7473 \text{ Å})$. The angles between these six bonds are 60°. According to the differences of the bondlength a distorted hexagon results. The trigobase is formed as a perfect equilateral triangle by three sulfur atoms with equal Tl-S distances of 3.3695 A. In the case of an ideal "hex-tri" antiprism the rotation-angle between the two bases is 30°. However, in erniggliite this angle was found to be 22°, indicating a twisting of the trigonal base for 8° towards the longer Tl-S bonds (Fig. 6a).

The mean value of all Tl-S distances for the "hex-tri" antiprism is 3.4616 Å, which is somewhat longer than mean values known from other Tl-sulfosalts with 9-fold coordinated thallium e.g. 3.38 Å in $T_{18}Pb_4Sb_{21}As_{19}S_{68}$ (the mean value of eight (Tl-S)-polyhedra; Nagl, 1979), 3.389 Â in synthetic parapierrotite (ENGEL, 1980), 3.369 Å in rebulite (BALIC- \angle UNIC et al., 1982), 3.43 Å in simonite (ENGEL et al., 1982). Except of simonite and rebulite all referenced (Tl-S)-polyhedra are distorted trigonal prisms with three additional sulfur atoms opposite to the three side faces.

The agreement of the mean value of the six shorter Tl-S bonds of 3.3188 À with the sum of the ionic radii, 3.32 Å of $T¹⁺$ and $S²⁻$ as found in TIS (Hahn and Klingler, 1949) and published to be r_{T1} + = 1.44 Å and r_s 2- = 1.84 Å (PAULING, 1968) suggest that the Tl-S bonding is mainly ionic. The three longer Tl-S bonds may indicate ^a chemical activity of the lonely electron pair of the thallium ion. In this structure ^a remarkable short Tl-As distance of 3.2594 À is observed which is slightly shorter than the shortest Tl-S bonding of 3.2681 Ä, whereas Tl-As interactions in other (Tl-As)-sulfosalts are usually in the range of 3.3 to 3.6 A. The stereochemical activity mentioned above and a possible Tl-As interaction may explain the shift of the Tl atom towards the hexagobase. Since in the structure of erniggliite the shortest Tl-Tl distance is 4.205 À we think there is no tendency of Tl-Tl interactions as proposed for other Tl-sulfosalts e.g. lorandite with Tl-Tl tances of 3.54 and 4.03 Å (FLEET, 1973).

The Sn-polyhedron is a nearly perfect trigonal antiprism with six equal Sn-S bonds of 2.5614 A and angles of 82.607°($3 \times$) and 97.393° ($3 \times$) (Figs 4, 6c). Inspecting the ICSD-file (BERGERHOF et al., 1983) we did not find any (Tl-Sn)-sulfosalt referenced up to now. So we searched for $(Tl-Sn)$ sulfides containing Sn with ^a formal valence of 4+, because we assume that in the structure of erniggliite this Sn-valence is present. We found four entries, two of them refer to structures with ^a more or less distorted tetrahedral surrounding (Klepp, 1984 and Piffard, 1984), in the third structure cited the T1 atom is coordinated by five sulfur in form of ^a distorted trigonal bipyramid (Eulenberger, 1981), whereas in the fourth (Tl-Sn)-sulfide additionally ^a deformed octahedron occurs (Piffard et al., 1984).

The As atom forms with three sulfur atoms a perfect trigonal pyramid with the As atom at the apex and with As-S distances of 2.2421 Å $(3 \times)$ and S-As-S angles of 96.691° (3 \times). This coordinapolyhedron is quite common in As-sulfosalts, but was not observed before in this perfect way with the exception of ellisite, $TI₃AS₃$ (Gostonc, 1980). The distances and angles in erniggliite are very close to the mean values for AsS_3 -pyramids in sulfosalts (2.29 Å, 98.0°, EDENHARTER, 1976).

The structure of erniggliite is based on layers // (0001) (Fig. 5). The layer consists of the TIS_{q} -" hex-tri"-antiprisms and isolated AsS_3 -pyramids, sharing common edges and corners. The thicknes of the layer corresponds exactly to the c-lattice parameter. The layers are connected only by the Sn antiprisms sharing common edges and corners with the polyhedra of the layers such that one half of the antiprism belongs to each layer respectively. This characteristic feature of the structure explains the very good cleavage of the crystals // (0001). The Tl-As interaction as mentioned above may additionally stabilize the connection of the layers.

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